



Originally published as:

Balasis, G., Egbert, G. D. (2006): Empirical orthogonal function analysis of magnetic observatory data: further evidence for non-axisymmetric magnetospheric sources for satellite induction studies. - *Geophysical Research Letters*, 33

DOI: [10.1029/2006GL025721](https://doi.org/10.1029/2006GL025721)



2 Empirical orthogonal function analysis of magnetic observatory data: 3 Further evidence for non-axisymmetric magnetospheric sources for 4 satellite induction studies

5 Georgios Balasis¹ and Gary D. Egbert²

6 Received 24 January 2006; revised 24 March 2006; accepted 8 May 2006; published XX Month 2006.

8 [1] Although satellite electromagnetic induction studies
9 have usually assumed a symmetric magnetospheric ring
10 current source, there is growing evidence for significant
11 source asymmetry. Here we apply empirical orthogonal
12 function methods to mid-latitude night-side hourly mean
13 geomagnetic observatory data to search for evidence of non-
14 zonal low-frequency source fields. The dominant spatial
15 mode of variability in residuals, obtained by subtracting
16 symmetric ring current and ionospheric fields of the CM4
17 comprehensive model, has a substantial Y_2^{-1} quadrupole
18 component and is highly correlated with D_{st} . This pattern of
19 temporal variability, which implies enhanced ring current
20 densities in the dusk sector, persists even when peak storm-
21 time data are omitted. The observed asymmetry agrees with
22 that inferred previously by Balasis et al. (2004), from the
23 local time dependence of biases in satellite induction
24 transfer functions. Temporal correlation of the leading
25 mode with D_{st} , and consistency of its spatial structure with
26 recent empirical ring current models, suggest a
27 magnetospheric origin. **Citation:** Balasis, G., and G. D.
28 Egbert (2006), Empirical orthogonal function analysis of
29 magnetic observatory data: Further evidence for non-
30 axisymmetric magnetospheric sources for satellite induction
31 studies, *Geophys. Res. Lett.*, 33, LXXXXX, doi:10.1029/
32 2006GL025721.

34 1. Introduction

35 [2] The traditional approach to estimation of the electrical
36 conductivity of Earth's mantle is based on interpretation of
37 ground-based observatory recordings of geomagnetic varia-
38 tions of external origin on time scales from hours to months
39 [e.g., Banks, 1969; Schultz and Larsen, 1990; Olsen, 1998;
40 Fujii and Schultz, 2002]. This approach has a serious
41 inherent limitation: the global distribution of magnetic
42 observatories is irregular and sparse, leaving large areas of
43 the Earth (especially the ocean basins) unsampled. Recent
44 magnetic satellite missions, such as Ørsted, CHAMP, and
45 SAC-C provide nearly complete global coverage, and thus
46 offer exciting possibilities for new insight into 3D patterns
47 of mantle conductivity. However, to date electromagnetic
48 (EM) induction studies with satellite data [e.g., Olsen, 1999;
49 Constable and Constable, 2004; Martinec and McCreddie,
50 2004] have all been based on very simple models of the

external sources: a symmetric magnetospheric ring current 51
(RC) described by Y_1^0 , and possibly a few other zonal 52
harmonics. 53

[3] Balasis et al. [2004] (hereinafter BEM) show that 54
estimates of EM induction transfer functions (TFs) obtained 55
from CHAMP data under the traditional assumption of a 56
symmetric RC source have biases which depend systemat- 57
ically on local time (LT). This pattern of biases suggests that 58
a purely zonal source model is inadequate. BEM further 59
showed that the pattern could be explained by adding a Y_2^1 60
quadrupole term correlated with the traditional axial dipole 61
source variations, and oriented so that meridional magnetic 62
fields peak in the dusk sector (at 19:30 LT). 63

[4] There has long been evidence for asymmetry in the 64
RC, particularly for the storm main phase [e.g., Daglis and 65
Kozyra, 2002; Daglis et al., 2003]. But are these asymme- 66
tries only brief, lasting for a few hours near storm onset? Or 67
are they persistent enough that they must be accounted even 68
for very long period induction studies, as the results of BEM 69
suggest? Only recently has clear evidence been presented 70
for asymmetry of the RC at all activity levels [e.g., 71
Jorgensen et al., 2004; Le et al., 2004]. These two studies 72
used in situ satellite data to directly map long term average 73
magnetospheric current densities as a function of D_{st} . 74
Systematic differences of average current densities between 75
activity levels, and persistence of asymmetries (with the 76
strongest fields centered in the midnight and dusk sectors) 77
suggest long period variations of the asymmetric RC, 78
closely coupled to activity level and D_{st} , consistent with 79
the indirect inference of asymmetry by BEM. 80

[5] Here we use geomagnetic observatory data to inves- 81
tigate this issue further, applying a novel analysis to 82
emphasize signals which vary slowly in a solar magnetic 83
(SM) reference frame. A major advantage of this data set is 84
that each observatory sweeps through all LTs once per day, 85
providing direct observation of non-zonal structure. A 86
disadvantage is that ground-based observations by them- 87
selves cannot distinguish magnetospheric and ionospheric 88
sources, and ionosphere currents exhibit very strong LT 89
dependence. We take two steps to minimize ionospheric 90
complications: (1) we subtract the CM4 comprehensive 91
model [Sabaka et al., 2004] ionospheric correction; and 92
(2) we focus on night-side mid-latitude data where iono- 93
spheric contamination is expected to be least. 94

2. Data Processing 95

[6] We analyzed three component hourly mean geomag- 96
netic data, starting with 79 observatories at all latitudes 97
(Figure 1), for 4 years (1997–2000). The following prelim- 98
inary processing steps were then applied: 99

¹GeoForschungsZentrum Potsdam, Potsdam, Germany.

²College of Oceanic and Atmospheric Sciences, Oregon State University, Corvallis, USA.

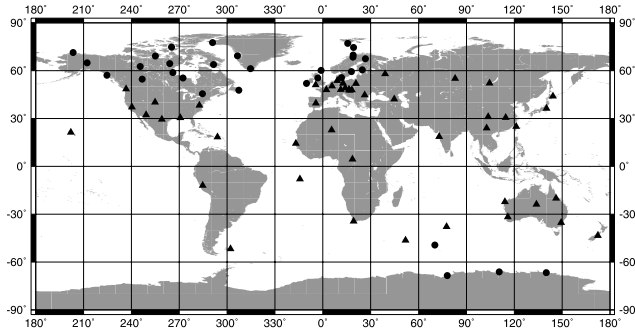


Figure 1. Geographic locations of geomagnetic observatories. Triangles denote mid-latitude observatories used for the EOF of Figure 3.

[7] 1) The CM4 model with the E-region ionospheric correction driven by the daily $F_{10.7}$ solar flux index, and with symmetric RC variations driven by D_{st} was subtracted. This removes much (but certainly not all) of the regular daily variation of ionospheric origin, approximately accounting for modulation in the size of the ionospheric cavity [Sabaka et al., 2004]. The D_{st} dependent part of CM4 correction also removes most of the axisymmetric RC variation and corresponding induced fields. After correction with the CM4, residual observatory means were subtracted (to remove short wavelength crustal fields not represented in the CM4), and each channel was high-pass filtered with a 50 day cutoff.

[8] 2) Time series for each observatory were interpolated from the standard hourly sampling at fixed UT, to fixed hourly sampling in magnetic local time (MLT). Daily sections of the filtered residual magnetic fields were then constructed with the data for all observatories aligned by MLT, and sorted by geomagnetic latitude (Figure 2). For the 4 year period this results in 1439 one day sections each with 79×24 “pixels”, each corresponding to one observatory at a fixed MLT. With this alignment, low frequency ($f \ll 1$ cycle per day) external source variations in a SM frame are sampled at all longitudes (and at the geomagnetic latitudes of the observatories) once per day.

[9] 3) Time series for each pixel were low-pass filtered with a 5 day cut-off. This filtering reduces the effects of more rapid variations which might be more appropriately described in UT, but does not smooth non-axisymmetric spatial structure in the SM frame. As illustrated in Figure 2, the combined effect of this filtering and alignment of the data by MLT produces smoother and cleaner looking daily samples of the magnetic fields.

[10] 4) Seasonal means, averaging over all days in the winter (Nov–Feb), summer (May–Aug) and equinox months were subtracted. These means are strongly dominated by auroral current systems (which CM4 does not explicitly model) but there are also spatially coherent corrections of the order of 5 nT at mid-latitudes.

[11] Special care was taken for missing data in steps 1–4: they were excluded from computation of means, set to 0 after subtracting means, and omitted from the calculation of averages when pixels were low-pass filtered. After these initial processing steps, we applied an empirical orthogonal function analysis (EOF, also known as principal component analysis) [see Preisendorfer, 1988] to the time sequence of

residual daily magnetic field vector variations. The EOFs decompose the total time varying signal into a sum over spatio-temporal modes: $H(\theta, \phi, t) = \sum_k X_k(\theta, \phi) T_k(t)$. Note that here θ represents observatory (geomagnetic) latitude, and ϕ MLT. The leading spatial mode $X_1(\theta, \phi)$ is the pattern of three component magnetic field vectors that explains the most variance in the sequence of daily sections. The corresponding temporal function $T_1(t)$ gives the time varying coefficient (or loading) of the spatial mode. Additional modes are orthogonal in space, and in time (i.e., the temporal modes are uncorrelated), and explain successively less of the residual variance. The EOF decomposition is essentially just the singular value decomposition (SVD) of the full data matrix, where rows of the data matrix represent spatial position (both latitude and longitude), and the columns represent replicates over time. The singular values give the relative amplitudes of each data mode. For the observatory analysis described here roughly 50% of the residual variance is contained in the first 3–4 modes with roughly 20% in the first mode alone, depending on details of the pre-processing.

3. Results

[12] The leading EOF modes obtained from analyzing observatories from all latitudes and all MLTs are dominated by auroral current systems, but there are significant large scale corrections at mid-latitudes as well. To focus on the asymmetries in magnetospheric sources suggested by the BEM results, which were based on mid-latitude night-side satellite data, we restricted the EOF analysis to observato-

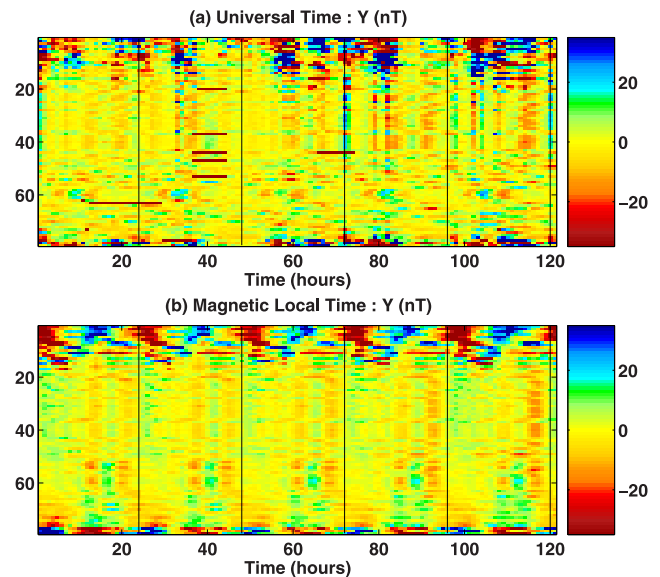


Figure 2. Example of residuals (geomagnetic east component) for 5 days of data from 79 observatories, after subtraction of CM4 predictions, (a) before and (b) after aligning and smoothing as described in text. The data are sorted by geomagnetic latitude N to S (axis on left is observatory number with geomagnetic dipole equator near number 65). Vertical lines denote day boundaries; each 79×24 panel is one (spatial) realization for the EOF decomposition.

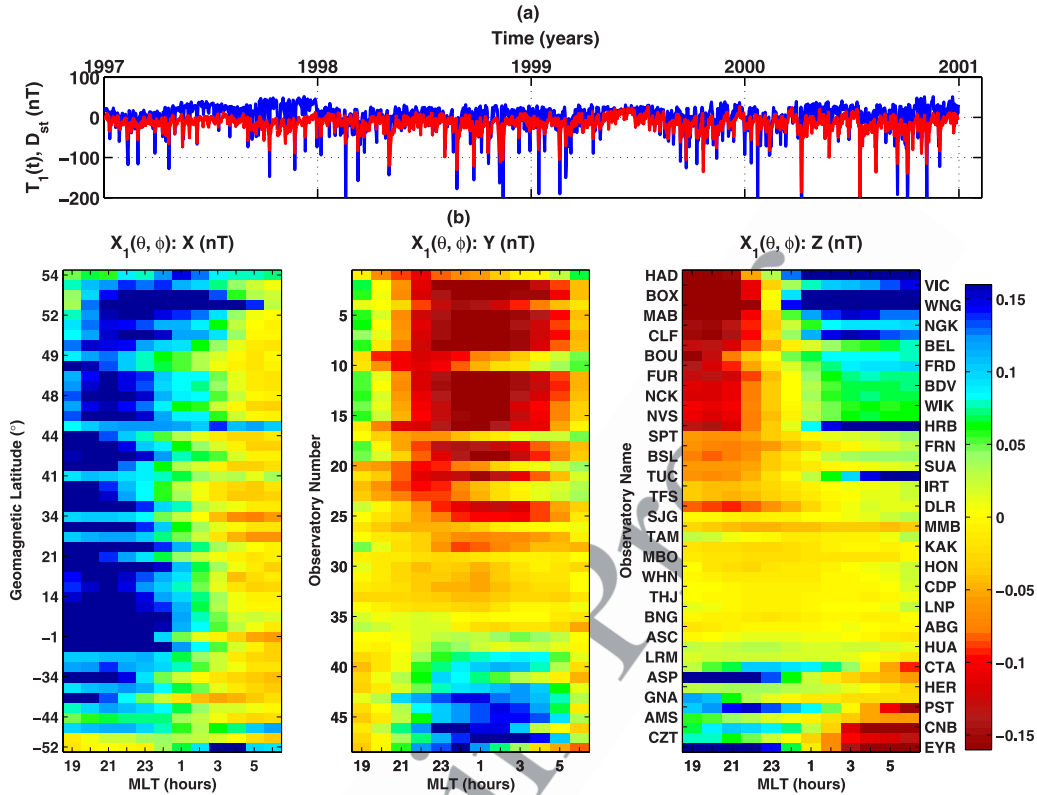


Figure 3. First EOF of the mid-latitude night-side magnetic observatory data, after processing outlined in text. (a) EOF time variation $T_1(t)$ (blue curve) scaled to be comparable to D_{st} (red curve). (b) Three components of the spatial mode $X_1(\theta, \phi)$, in geomagnetic coordinates. Geomagnetic latitude, observatory number, and code are given on the y axes of the X, Y and Z plots, respectively.

175 ries between 50°S and 50°N geomagnetic, and to MLTs for
 176 the 12 hour period centered on local midnight. The first
 177 mode (Figure 3) indeed shows clear evidence for the large
 178 scale non-axisymmetric signal inferred by BEM. Note that
 179 the temporal variation of this signal is very clearly correlated
 180 with D_{st} (Figure 3a). The positive correlation of $T_1(t)$ with
 181 D_{st} means that fluctuations in magnetic fields on the Earth
 182 surface associated with the storm time variations tracked by
 183 D_{st} exhibit persistent asymmetries, with peak amplitudes
 184 centered in the dusk sector. In Figure 3 $T_1(t)$ is scaled to be
 185 comparable to D_{st} , so the scale of the spatial mode gives the
 186 relative amplitude, normalized to the symmetric RC com-
 187 ponent, of the asymmetry. The first EOF thus corresponds to
 188 variations in the north component (X) that are of the order of
 189 10–15% of the symmetric D_{st} component. The other leading
 190 EOFs (not shown) also exhibit large scale spatial structure
 191 but are increasingly noisy. These EOFs also reveal evidence
 192 for seasonal variations, suggesting that the treatment of
 193 seasonal effects in CM4 could be further refined. When
 194 storm commencement data are excluded prior to step 3 (e.g.,
 195 omitting all data with $D_{st} < -75$ nT) very similar results are
 196 obtained for the EOF analysis.

197 [13] We used least squares to fit the horizontal compo-
 198 nents of the leading spatial mode ($12 \times 48 \times 2 = 1152$
 199 elements) as the gradient of a scalar potential expanded in
 200 spherical harmonics to degree and order 5 (35 parameters).
 201 A backward elimination procedure, using a 0.99 signifi-
 202 cance level [e.g., Kleinbaum and Kupper, 1978], was used
 203 to eliminate spherical harmonics which did not contribute

significantly to the fit. The resulting model includes 5 terms
 and fits the EOF with $R^2 = 0.82$, nearly as well as the full
 degree 5 expansion ($R^2 = 0.85$). Three terms dominate:
 together Y_2^{-1} , Y_1^0 , and Y_4^{-1} explain most of the variance ($R^2 =$
 0.79).

4. Discussion

[14] Empirical orthogonal function analysis of night-side
 (18:00–06:00 MLT) mid-latitude (50°S to 50°N) observa-
 tory data show clear evidence for large scale non-axisym-
 metric structure that is fixed in MLT, coherent with D_{st} , and
 persists even when storm commencement data are excluded.
 The significant Y_2^{-1} quadrupole component in the first mode
 EOF magnetic fields implies an enhancement of RC density
 in the dusk sector, and meridional current on the night-side
 centered near local midnight. Examination of Figure 3
 suggests that the peak in RC density is in fact shifted
 several hours toward local midnight, consistent with the
 alignment of the asymmetry inferred by BEM. The pattern
 in the dominant EOF is also in broad agreement with the
 recent empirical RC models of Jorgensen *et al.* [2004] and
 Le *et al.* [2004], which exhibit persistent peaks in outer RC
 density in the dusk and midnight sectors. In particular Le *et al.*
 found that on average the peak of the RC shifts toward
 the dusk sector for higher D_{st} levels. This is consistent with
 the strong positive correlation between D_{st} and the dominant
 EOF from our analysis. The spatial structure of the domi-

- 230 nant EOF also agrees well with the *Tsyganenko* [2002] 288
 231 model for the partial RC. 289
 232 [15] We cannot completely rule out the possibility that the 290
 233 asymmetry observed in the observatory data has an iono- 291
 234 spheric origin. However, in addition to the agreement 292
 235 between our results and the empirical maps of magneto- 293
 236 spheric RC densities, two further lines of evidence support a 294
 237 magnetospheric source for the observed asymmetries. Most 295
 238 striking is the very strong correlation of the temporal 296
 239 loading of the dominant EOF with D_{st} (Figure 3a). This 297
 240 strong correlation would be very surprising if the asymme- 298
 241 try were not due dominantly to magnetospheric sources 299
 242 [Daglis and Kozyra, 2002]. Second, BEM indirectly in- 300
 243 ferred the same pattern of magnetospheric RC asymmetry
 244 from the LT dependence of induction TF biases estimated
 245 from CHAMP satellite magnetic data. Since CHAMP flies
 246 above the ionosphere and below the magnetosphere, the
 247 satellite data can distinguish between ionospheric and
 248 magnetospheric sources. BEM found that the observed
 249 pattern of biases could be readily explained by addition of
 250 a quadrupolar source peaked near dusk (i.e., essentially Y_2^{-1})
 251 that was external to the satellite orbit (i.e., magnetospheric),
 252 but not by similar internal (i.e., ionospheric) non-axisym-
 253 metric sources.
- 254 [16] Because of Earth rotation any slowly varying non-
 255 axisymmetric structure will in fact result in induction at
 256 daily variation periods; only the axisymmetric part of the
 257 RC will contribute to induction at long periods. However,
 258 accurate models of the non-axisymmetric signal will still be
 259 essential to proper interpretation of the satellite data, since
 260 these components will appear in the satellite frame as
 261 slowly varying components of the magnetic field. Given
 262 the growing body of evidence for asymmetry in the RC, we
 263 suggest that further progress in satellite induction studies
 264 will require moving beyond the simple axisymmetric source
 265 field models used to date. Proper interpretation of the
 266 satellite data will also require accurate separation of iono-
 267 spheric and magnetospheric sources, as these will be re-
 268 spectively internal and external to the satellite orbit.
 269 Interconnections between magnetospheric and ionospheric
 270 current systems will also have to be modeled properly. For
 271 example, the meridional currents associated with the domi-
 272 nant EOF of Figure 3 almost certainly flow along field
 273 lines, and close in the auroral ionosphere. Thus, these
 274 currents will be partly external, and partly internal, to the
 275 satellite orbit.
- 276 [17] Combination of both satellite and observatory data
 277 will be essential to the task of developing improved source
 278 models for induction studies. Our study here using only
 279 observatory data, combined with the study of BEM using
 280 only satellite data provides a glimpse of the potential power
 281 of combining these multiple data sources. The asymmetry is
 282 well mapped by the observatories in this study; incorporat-
 283 ing the CHAMP results from BEM implies that the asym-
 284 metry most likely comes from the magnetosphere. Future
 285 work should combine these two data sources more explicitly
 286 to further improve separation and characterization of iono-
 287 spheric and magnetospheric sources. Obviously we also
- need to take advantage of the existing extensive observa-
 tional, theoretical, and modeling studies of the magneto-
 sphere to guide development of the more realistic models
 for magnetospheric (and ionospheric) sources that will be
 required for progress in satellite induction studies. Statistical
 estimates of average magnetospheric RC densities and/or
 semi-empirical models such as that of *Tsyganenko* [2002]
 might be useful starting points for these developments.
 Ultimately, data assimilation methods which combine phys-
 ics-based models of the magnetosphere and ionosphere with
 all available data offer the greatest hope for accurate
 modeling of external sources for global induction studies
 with satellite data.
- [18] **Acknowledgment.** Constructive suggestions from two anony-
 mous reviewers are gratefully acknowledged.
- ## References
- Balasis, G., G. D. Egbert, and S. Maus (2004), Local time effects in satellite
 estimates of electromagnetic induction transfer functions, *Geophys. Res.
 Lett.*, *31*, L16610, doi:10.1029/2004GL020147.
- Banks, R. J. (1969), Geomagnetic variations and the conductivity of the
 upper mantle, *Geophys. J. R. Astron. Soc.*, *17*, 457–487.
- Constable, S., and C. Constable (2004), Observing geomagnetic induction
 in magnetic satellite measurements and associated implications for mantle
 conductivity, *Geochem. Geophys. Geosyst.*, *5*, Q01006, doi:10.1029/
 2003GC000634.
- Daglis, I. A., and J. U. Kozyra (2002), Outstanding issues of ring current
 dynamics, *J. Atmos. Sol. Terr. Phys.*, *64*, 253–264.
- Daglis, I. A., J. U. Kozyra, Y. Kamide, D. Vassiliadis, A. S. Sharma, M. W.
 Liemohn, W. D. Gonzalez, B. T. Tsurutani, and G. Lu (2003), Intense
 space storms: Critical issues and open disputes, *J. Geophys. Res.*, *108*,
 1208, doi:10.1029/2002JA009722.
- Fujii, I., and A. Schultz (2002), The 3D electromagnetic response of the
 Earth to ring current and auroral oval excitation, *Geophys. J. Int.*, *151*,
 689–709.
- Jorgensen, A. M., H. E. Spence, W. J. Hughes, and H. J. Singer (2004), A
 statistical study of the global structure of the ring current, *J. Geophys.
 Res.*, *109*, A12204, doi:10.1029/2003JA010090.
- Kleinbaum, D. G., and L. L. Kupper (1978), *Applied Regression Analysis
 and Other Multivariable Methods*, Duxbury, Pacific Grove, Calif.
- Le, G., C. T. Russell, and K. Takahashi (2004), Morphology of the ring
 current derived from magnetic field observations, *Ann. Geophys.*, *22*,
 1267–1295.
- Martinez, Z., and H. McCreadie (2004), Electromagnetic induction model-
 ing based on satellite magnetic vector data, *Geophys. J. Int.*, *157*, 1045–
 1060.
- Olsen, N. (1998), The electrical conductivity of the mantle beneath Europe
 derived from *C*-responses from 3 to 720 hr, *Geophys. J. Int.*, *133*, 298–
 308.
- Olsen, N. (1999), Induction studies with satellite data, *Surv. Geophys.*, *20*,
 309–340.
- Preisendorfer, R. W. (1988), *Principal Component Analysis in Meteorology
 and Oceanography*, 425 pp., Elsevier, New York.
- Sabaka, T. J., N. Olsen, and M. E. Purucker (2004), Extending compre-
 hensive models of the Earth's magnetic field with Oersted and CHAMP data,
Geophys. J. Int., *159*, 521–547.
- Schultz, A., and J. C. Larsen (1990), On the electrical conductivity of the
 mid-mantle, II, Delineation of heterogeneity by application of external
 inverse solutions, *Geophys. J. Int.*, *101*, 565–580.
- Tsyganenko, N. A. (2002), A model of the near magnetosphere with a
 dawn-dusk asymmetry 2. Parameterization and fitting to observations,
J. Geophys. Res., *107*, doi:10.1029/2001JA000220.
- G. Balasis, GeoForschungsZentrum Potsdam, Telegrafenberg, D-14473
 Potsdam, Germany. (gbalasis@gfz-potsdam.de)
 G. D. Egbert, College of Oceanic and Atmospheric Sciences, Oregon
 State University, Oceanography Admin. Bldg. 104, Corvallis, OR 97331–
 5503, USA. (egbert@coas.oregonstate.edu)